



Exploring the concept of assigning relative weights to regional climate models: Experiences from the ENSEMBLES project

Acknowledgements to ENSEMBLES RT3 Partners (ICTP,CNRM,CUNI,DMI,HC,KNMI,MPI,SMHI,UCLM)

Erik Kjellström Presentation at the Workshop on Uncertainties of scenario simulations, SMHI 14 October 2010

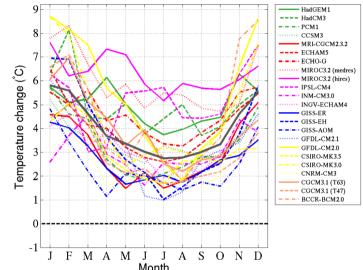




Multi-model ensemble predictions/projections can be characterized by a large uncertainty due to the inter-model spread



Climate change in Northern Sweden Comparing 2071-2100 vs 1961-1990 All models run under SRES A1B



- Different models within the ensemble might have different levels of performance in climate simulation/prediction
- By weighting the models based on their "performance" it might be possible to reduce the uncertainty and increase the reliability of the prediction/projection





	Global model Regional inst.	METO-HC	METO-HC Low sens.		MPIMET Standard	MPIMET Ens.m. 1	MPIMET Ens.m. 2	IPSL	CNRM	NERSC	MIROC	сссмз	Total number
	МЕТО-НС	2100	2100*	2100*	2100 (late 2010)								4
(RCMs)	MPIMET				2100			2050*					2
Ö	CNRM								2100				1
<u>R</u>	DMI				2100*				2100	2100* (04/2010)			3
<u>S</u>	ETH	2100											1
models	кимі				2100* 2100	<u>2100</u> *	2100*				<u>2100</u> *		1+4
Ĕ	ІСТР				2100								1
ate	SMHI		2100*		2100* 2100*					2100			3+1
climate	UCLM	2050											1
<u>U</u>	C4I			2100*		2050 (A2)*							2
Ø	GKSS							2050*					1
0	METNO	2050*								2050*			1
egional	СНМІ								2050* (12/2009)				1
Ř	OURANOS**											2050*	1
	VMGO**	2050*											1
	Total (1951- 2050)	5	2	2	7+2	0+1	0+1	2	3	3	0+1	1	25+5

Global climate models (GCMs)

The ENSEMBLES GCM-RCM Matrix

Red: Online now; *: non-contractual runs; **:affiliated partners without obligations; <u>underscore</u>: 50km resolution; (in parantheses): Expected date. For partner acronym explanations, see the participant list. **NOTE** that all partners also did an ERA-40 driven analysis 1951(1961)-2000

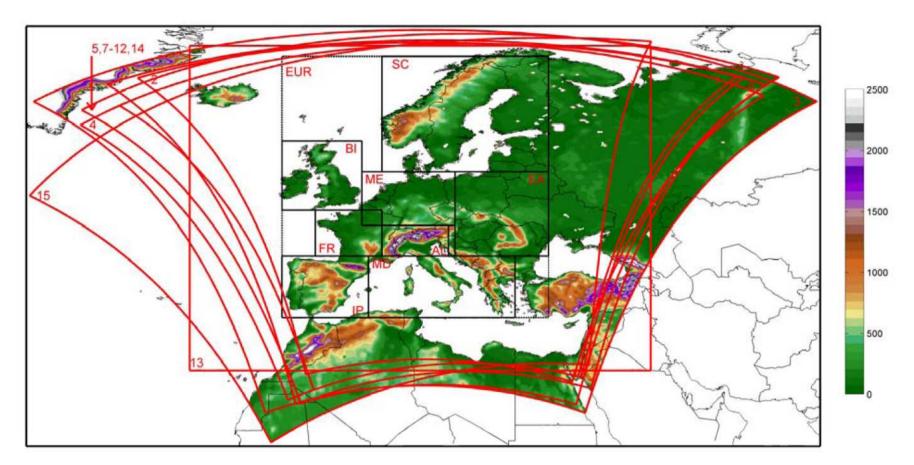




- Develop weights based on different metrics of model performance in reproducing present day climate characteristics, with emphasis on the "added value" obtained from RCMs
- Six metrics were identified (based on ERA40-driven runs)
 - F1: Large scale circulation and weather regimes (CNRM)
 - F2: Temperature and precipitation meso-scale signal (ICTP)
 - F3: PDFs of daily precipitation and temperature (DMI, UCLM,SHMI)
 - F4: Temperature and precipitation extremes (KNMI; HC)
 - <u>F5: Temperature trends (MPI)</u>
 - F6: Temperature and precipitation annual cycle (CUNI)
- Weights have been calculated for single seasons and regions and then averaged to yield one final number per model







- 15 RCMs at 25km, lateral boundary conditions from ERA40
- Analysis period 1961-2000
- Common minimum domain, all data regridded to a common 25 km lat-lon grid
- Observations on monthly {CRU 0.16 degree (Mitchell et al. 2003)} and daily {EOBS 0.25 degree (Haylock et al. 2008)}

Christensen et al., Clim. Res. accepted for publication





- Mean behavior
 - Frequency of occurrence, W(1,1,2)
 - Spatial composite, W(1,1,1)
 - Duration, W(1,1,3)
- Interannual variability
 - Variance of the frequency of occurrence, W(1,2,1)
 - Temporal correlation, w(1,2,2)
- **Daily behavior**
 - Total number of days per season, W(1,3,1)

$$w_{1}^{i,s} = w_{111}^{\alpha_{111}} * w_{112}^{\alpha_{112}} * w_{113}^{\alpha_{113}} * w_{121}^{\alpha_{121}} * w_{122}^{\alpha_{122}} * w_{131}^{\alpha_{131}}$$
$$\widetilde{w}_{1}^{i,s} = \frac{w_{1}^{i,s}}{\sum_{i} w_{1}^{i,s}}$$
Samuel Somot (CNRM)



Weather regimes

30

20

10

0

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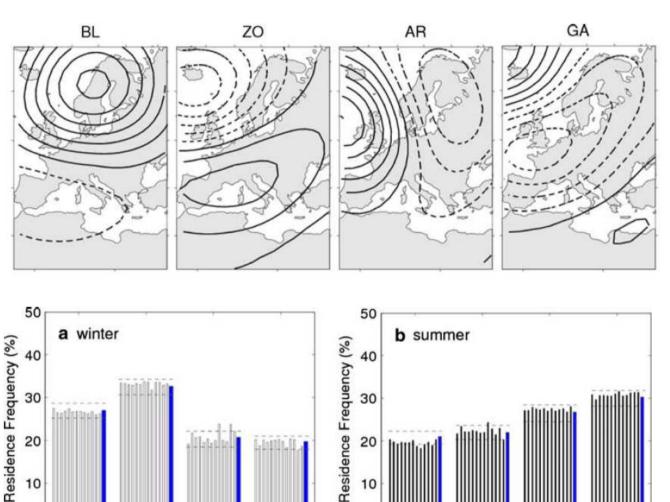
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- Daily Z500 data
- Clustering by PCA
- 4 regimes: Blocking Zonal Atlantic Ridge Greenland anticyclone



30

20

10

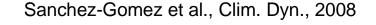
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- First run a smoother on the original fields to identify a large scale signal
- Define the mesoscale signal as the difference between the original fields and the large scale fields
- Define the 5 functions:

$$g_{1} = R(p)^{n} \qquad g_{2} = R(T)^{n} \qquad g_{3} = \sigma(t)^{p}_{CRU} / RMSE(p)$$
$$g_{4} = \sigma(t)^{T}_{CRU} / RMSE(T) \qquad g_{5} = \left[L - \left(\left| R_{pT}^{CRU} - R_{pT}^{mod} \right| / 2 \right) \right]$$

The weight is given by

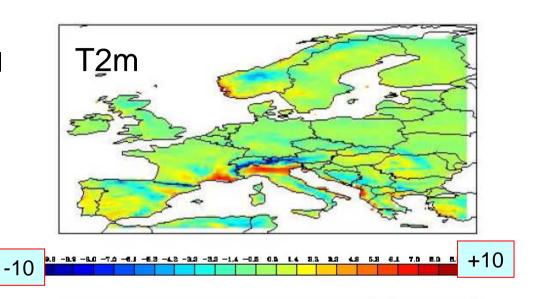
$$w_i = g_1 * g_2 * g_3 * g_4 * g_5$$

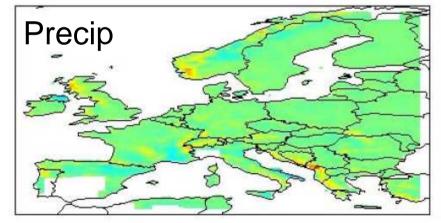


What is the mesoscale signal?



- Calculate 9x9 gridpoint spatial mean to get "large-scale" signal.
- Subtract the "large-scale" signal from the total field to get the "mesoscale" signal
- Particularly orographic features stands out. But also some coastal areas and large lakes





-20 -26 -28 -35 -84 -81 -17 -13 -29 -08 -08 08 08 0.9 13 17 21 34 85 83 28 29





F3: PDFs of daily and monthly temperature and precipitation



The skill score metric for daily data

Cumulative minimum of two distributions.

$$S_{\text{score}} = \sum_{1}^{n} \operatorname{minimum}(Z_m, Z_o),$$

Takes a value between 0 and 1

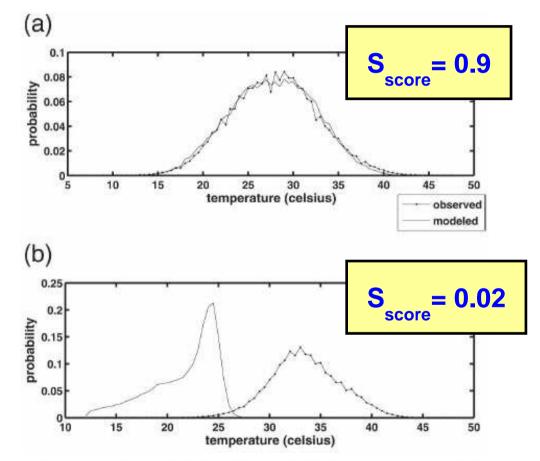


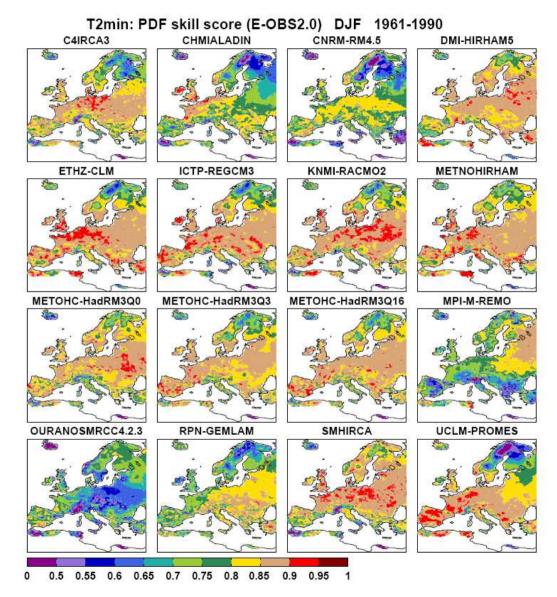
FIG. 3. Diagrams of modeled vs observed PDF illustrating the total skill score in (a) a near-perfect skill score test (0.9) and (b) a very poor skill score (0.02).

Perkins et al., J. Clim., 20, 2007



Skill scores based on daily data for winter ${\rm T}_{\rm min}$





Kjellström et al., Clim. Res. Accepted for publication



An alternative skill score metric for comparing CDFs



$$f_{1j} = 1 - \left(\frac{|A_{RCM_j} - A_{CRU}|}{2 \cdot A_{CRU}}\right)^{0.5}$$

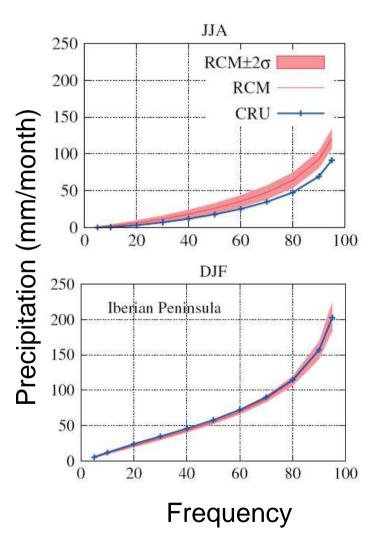
$$f_{2j} = 1 - \left(\frac{|A_{RCM_j}^+ - A_{CRU}^+|}{2 \cdot A_{CRU}^+}\right)^{0.5}$$

$$f_{3j} = 1 - \left(\frac{|A_{RCM_j}^- - A_{CRU}^-|}{2 \cdot A_{CRU}^-}\right)^{0.5}$$

$$f_{4j} = 1 - \left(\frac{|\overline{P_{RCM_j}} - \overline{P_{CRU}}|}{2 \cdot \overline{P_{CRU}}}\right)^{0.5}$$

$$f_{5j} = 1 - \left(\frac{|\sigma_{RCM_j} - \sigma_{CRU}|}{2 \cdot \sigma_{CRU}}\right)^{0.5}$$

 $W_j = f_{1j} \cdot f_{2j} \cdot f_{3j} \cdot f_{4j} \cdot f_{5j}$



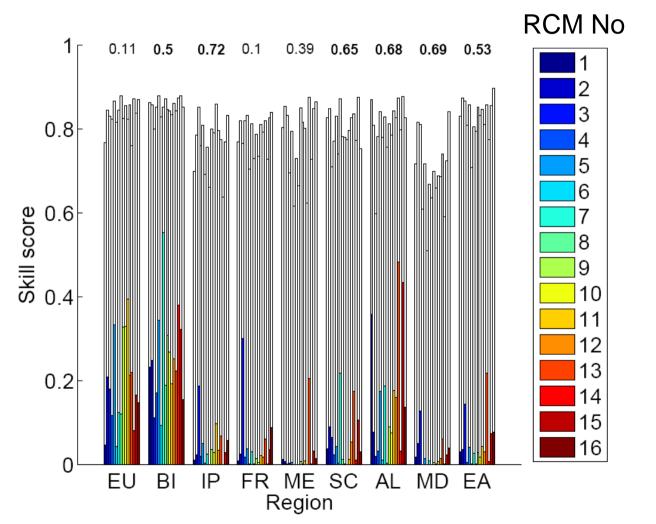
Sanchez et al., ASL, 2009



Comparing two different metrics for calculating skill scores



- Perkins (cumulative minimum) unfilled bars
- Sanchez (5 aspects of CDF match) – colored bars
- Absolute numbers are very different
- Ordering of RCMs differ



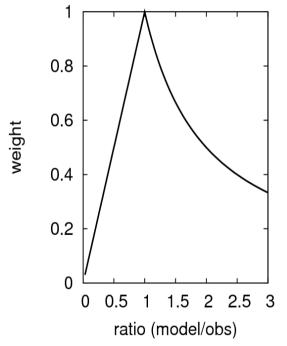




- Biases in extreme percentiles for temperature and precipitation are first calculated (99., 99.9, 99.99, 99.999 %)
- Biases (B) are then turned into weights (W) using an asymmetric transfer function

$$B = 100 \left[\frac{P_{rcm}}{P_{obs}} - 1 \right]$$

$$W = 1 + B/100,$$
 $B < 0$
 $= \frac{1}{1 + B/100},$ $B > 0$

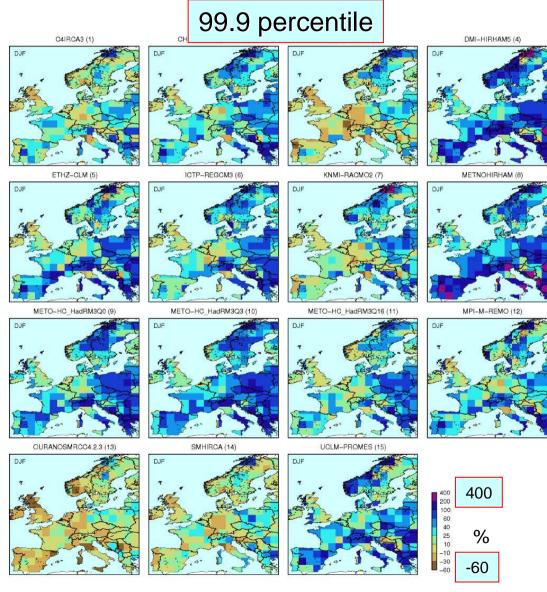


Lenderink., Clim. Res. accepted for publication



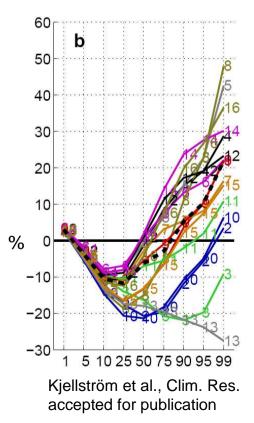
Biases in DJF extreme precipitation





Lenderink., Clim. Res. accepted for publication

- Wet biases compared to E-OBS in most but not all RCMs
- The spread between RCMs grow the further to the "wet side" one looks







 Linear temperature trends for the period 1961-2000 are calculated for observations and simulations

$$y_i = \alpha + \beta t_i + r_i$$

The trends are then turned into skill scores (or weights) using the formula

$$S = 1 - \frac{|\beta - \beta_{REF}|}{\zeta + |\beta - \beta_{REF}|}$$

 $\boldsymbol{\zeta}$ is a scaling parameter that determines the spread between the best/worst model

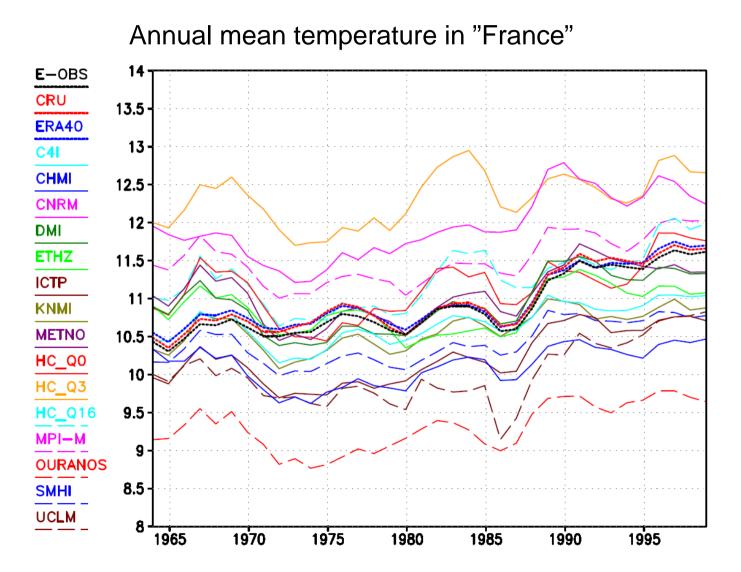
Annual and seasonal values are combined using

$$S_{combined} = 0.5 * S_{Year} + 0.125 * (S_{DJF} + S_{JJA} + S_{MAM} + S_{SON})$$

Lorenz and Jacob, Clim. Res. submitted







Lorenz and Jacob, Clim. Res. submitted

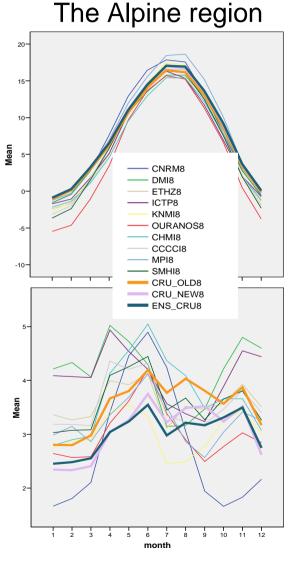




The weight depends on the model skill in reproducing amplitude and phase of the annual cycle

$$S = \frac{4(1+R)}{(\sigma + \frac{1}{\sigma})^{2}(1+R_{0})},$$

R=Correlation (Phase) σ=Ratio simulated/observed STD (magnitude)



Halenka et al., Clim. Res. submitted



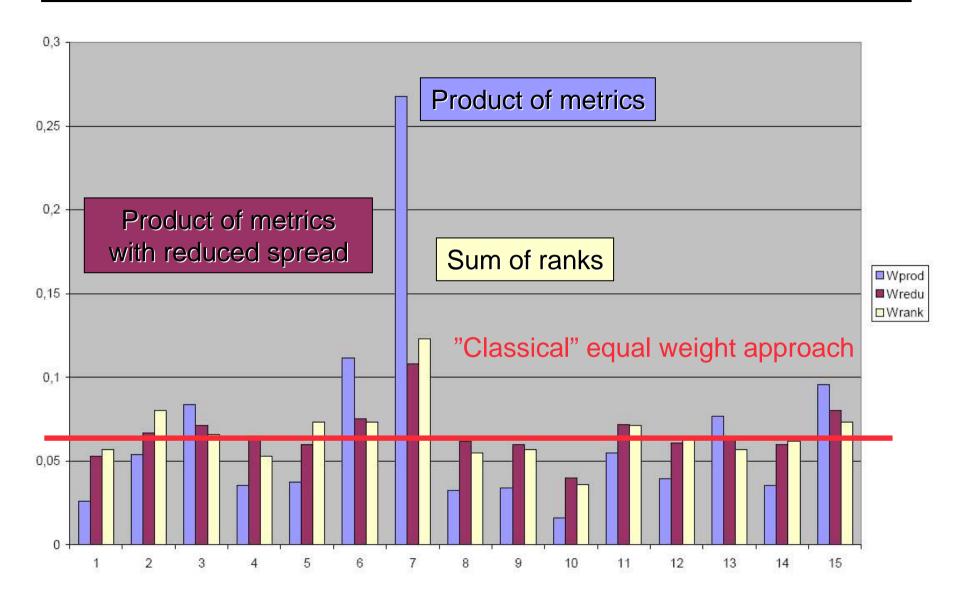


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- These were simply multiplied to yield the final weight $W_{RCM} = \prod_{i=1}^{6} f_i^{n_i}$
- A sensitivity study with a reduction in spread was performed (all single metric F1-F6 were allowed to be equally important)
- An alternative approach was to base the weight on ranking in the different metrics (F1-F6)



Final weights





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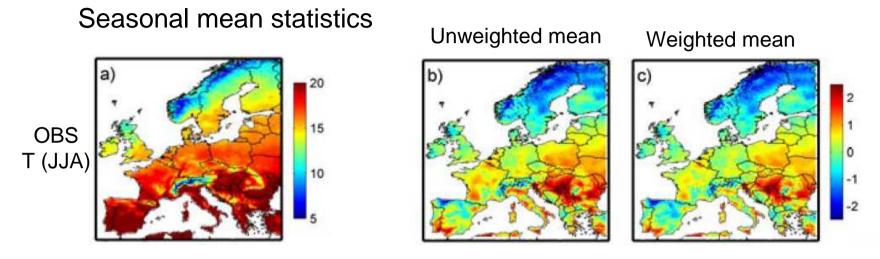
- Skill scores differ between different RCMs for different regions, seasons and variables
- Multiplicative metrics implies that the overall weight can be dominated by one "outlier" metric. One may overcome this problem by
 - Combine the metrics in different ways (addition rather than multiplication)
 - Adopt different metrics for different variables
 - Weight differently each metric
 - Normalize the metric by the inter-model spread
- Large subjective component of the approach
- The derived weights are to be used for the whole ensemble as they are derived relative to each other
- May be necessary to calculate other weights for certain impact studies
- When run with LBCs from GCMs also the GCMs could/should be weighted

A number of papers describing this work will appear in a special issue in Climate Research in late 2010 or early 2011



Does it matter?





(MAE – mean absolute error, RMSE – root mean square error, Areal fraction where MAE decreases)

	on	Areal fraction	RMSE	MAE	Variable
-		0.55	0.565/0.543	0.292/0.286	P (JJA)
ndicates	Bold face i	0.53	0.455 /0.465	0.372/0.377	P (DJF)
	improveme	0.74	0.928 /1.018	0.740/0.824	$T_{2m}\left(JJA\right)$
	 compared unweighted 	0.34	1.452/1.407	1.049/0.985	$T_{2m}\left(DJF\right)$
		0.63	12.87 /13.33	10.60 /11.10	$N_{\text{tot}}\left(JJA\right)$
-		0.45	7.97/7.70	7.34/6.08	N _{tot} (DJF)
		Christenson et a			

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